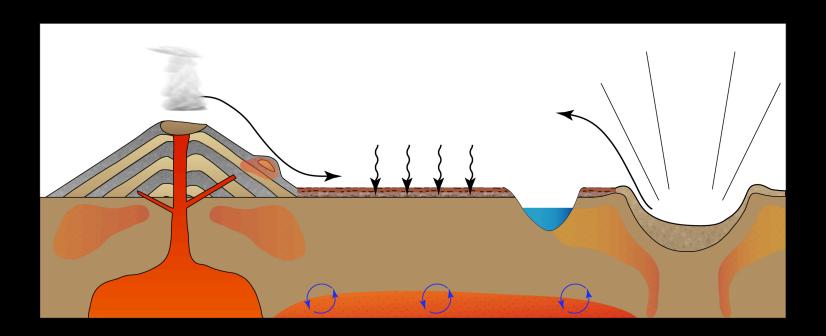
The early supracrustal rocks of Mars



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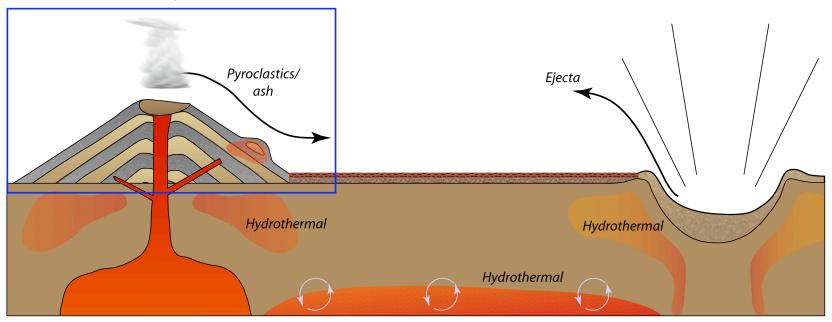
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Surface processes on Early (Noachian) Mars

Explosive volcanism / pyroclastic deposition



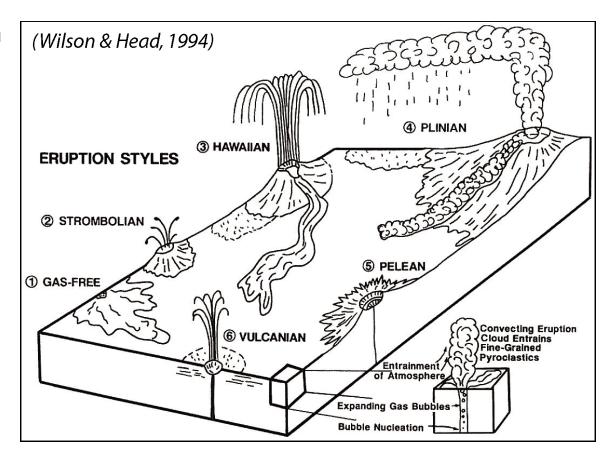
In the absence of geomorphic evidence for abundant surface water pre- 3.7 Ga we assume three processes dominate the earliest martian surface:

- 1. Impacts
- 2. Volcanism
- 3. Eolian reworking

Here, we explore rock accumulation and mineralogical predictions in this context Finally, we compare this model to observations at Mawrth Vallis

Explosive volcanism: An dominant contributor to the martian surface

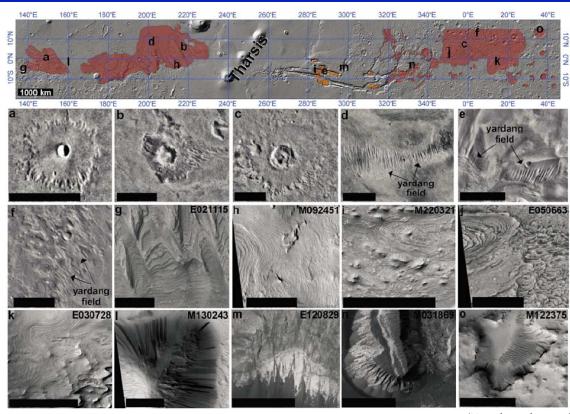
- Explosive volcanism on Mars common:
 - Lower atm pressure
 - Greater gas exsolution
 - Enhanced magma disruption
 - Higher volatile content
- Pyroclastic deposits
 - 100x finer grained
 - Greater distances
 - Clouds rise 5x higher



- Virtually every eruption (regardless of composition) of volatile-bearing magma accompanied by fine-grained pyroclastic deposit
- The Earth's sedimentary budget contains a minimum of 30% volcaniclastic contributions (Garrels & Mackenzie, 1971)

Explosive volcanism on Mars: Observational support

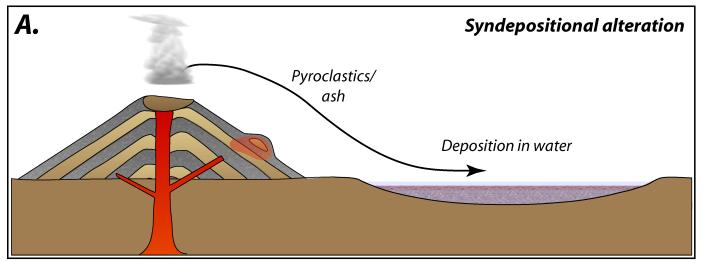
- Volcanic ash deposition important in forming Friable layered deposits (FLDs)
 - Mapped by Schultz & Lutz (1988) & Hynek et al. (2003) across equatorial regions 1000s km apart
- FLDs fine-grained & thinly layered
 - Distributed over 10⁵ to 10⁶ km²
 - Mostly post-Noachian



(Hynek et al., 2003)

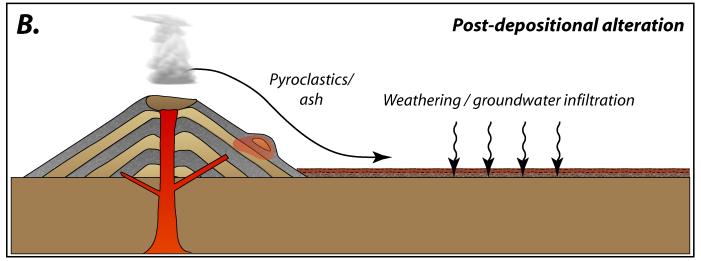
- Additional support for pyroclastic activity from Home Plate in Gusev Crater
- Many additional reports (Mouginis-Mark, 1982, 2002; Edgett, 1997; Greeley & Crown, 1990, etc.)
- Robbins et al. (2010) identify transition from explosive to effusive volcanism at about 3.5 Ga from survey of major calderas

Usually, we associate the products of explosive volcanism with either:



Requirements:

- Nearly pure smectite deposits (e.g., bentonites)
- Pre-existing water body must be spatially and temporally persistent
- Additional detrital pulses/signals?



Requirements:

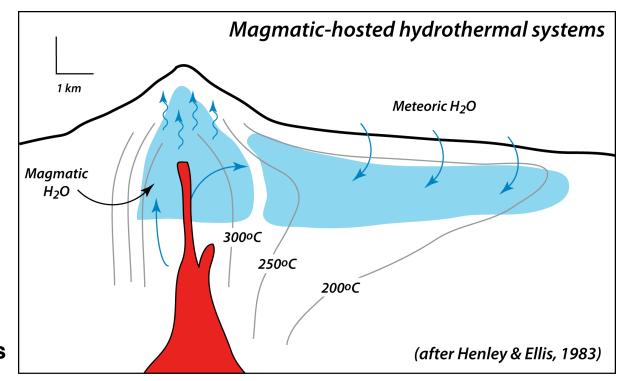
- Dominated by halloysite, kaolinite, opaline SiO₂, imogolite
- Smectite rare; when present is Al-rich

The origin of smectite in pyroclastic deposits

- Most smectite in pyroclastic deposits & soils is hydrothermally formed within volcanic vents and entrained upon eruption
 - Mineralogy, geochemistry, $\delta^{18}O$ & δD show elevated formation temperatures
 - Smectite is Fe/Mg-rich; including saponite & nontronite
 - Occurs with minor/trace chlorite, biotite, vermiculite, and C/S, etc.
 - Hydrothermal smectite deposited w/ash in localities across globe
- Mt. St Helens 1980 deposits (Pevear et al., 1982)
 - Largely andesitic/basaltic composition
 - Saponite in the unaltered ash reaches 10 vol. % of bulk, increasing with distance from vent
 - Total deposits >25% lithic material (i.e., entrained vent rocks)
- Mt. Usu, Japan (Mizota & Faure, 1998)
 - Nontronite present at >10 vol. % in unaltered ash
 - Water-saturated zone under vents driving hydrothermal activity

Hydrous magmatism & hydrothermalism: Hand in hand

- Every group of meteorites preserves evidence for magmatic hydrothermalism
- Abundant geochemical evidence for early hydrous volcanism
- When magmatic volatiles were high, degassing & hydrothermalism results



- Petrogenetic studies infer the earliest magmas were also the most volatile rich
- Thus, magmatic hydrothermal activity can be inferred to be directly proportional to volcanism
- Meteoric water may also contribute, but constraints are less clear pre-3.5 Ga

The pyroclastic contribution to early martian sedimentary budget

- Significant volcanism in pre-/mid-/late-Noachian (e.g., Tharsis)
- Magma-water/ice interactions would have produced:

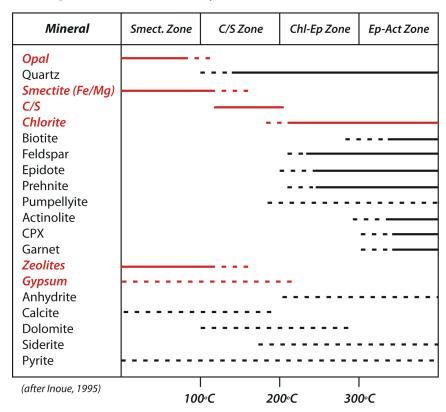
Hydrothermal systems

- Dependent on nature of interaction,
 volatile content, crustal H₂O reservoir
- Generate significant Fe/Mg-smectite & lesser chlorite, C/S, biotite, etc.
- Clays would be transported & concentrated in air fall deposits

Fine-grained pyroclastics

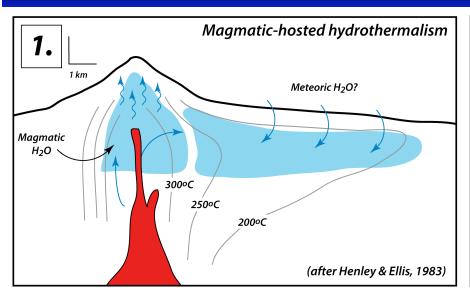
- Entrainment of atmosphere could alter syn-eruptive ash
- Distinct post-eruptive Al-rich weathering signal

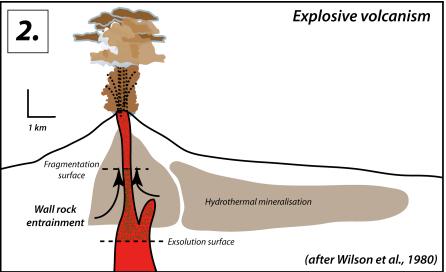
Ca/Mg-Series Alkaline Hydrothermal Mineralization

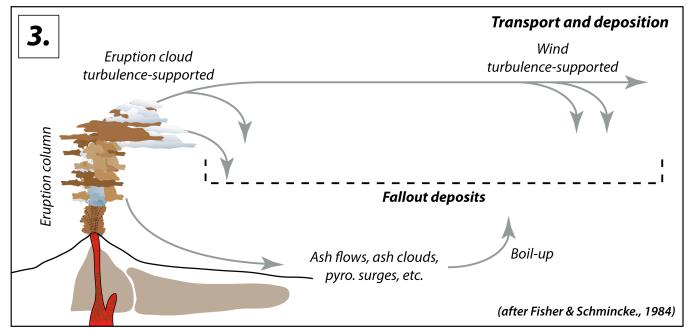


- The total amount of pyroclastic clay delivery to surface is a function of:
 - Frequency & extent of magmatic-hydrothermal & syn-eruptive alteration (dependent on volatiles)
 - Efficiency of explosive delivery & transport
- Volumetrically significant flux of newly formed clay available for impact/physical reworking

A model for pyroclastic contributions to surface mineralogy





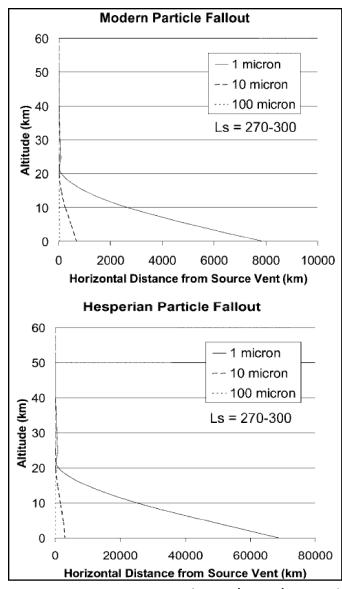


How much & how far? Ash transport and particle size segregation

 Atmospheric winds could carry <10μm particles several 1000 km

 Ordovician ash beds in E. US distributed over 5 x 10⁶ km²

- Significant atmospheric residence time (up to 4 years on Earth)
- Smectite & other clays present in ash concentrated in the <2µm fraction w/ increasing transport distance
 - e.g., Mt. St. Helens ash deposits as a function of distance from source



(Hynek et al., 2003)

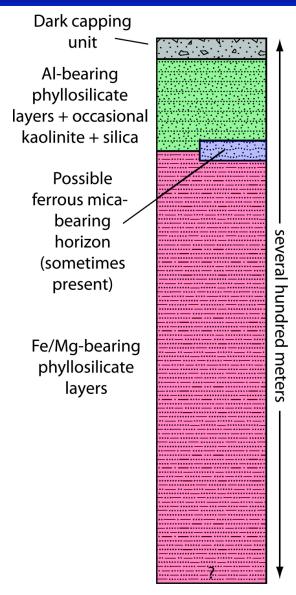
Implications for "layered" phyllosilicates (e.g., Mawrth Vallis)

Mawrth Vallis:

- Aerially extensive (>10⁶ km²)
- Mineralogy consistent within stratigraphic units
- Temporally significant
 - · Buried impact craters
- Compositional stratigraphy observed elsewhere
 - Eridania, Noachis, Valles Marineris

A significant pyroclastic contribution is capable of explaining:

- Persistent & extensive distribution
- Disparity between TES & OMEGA/CRISM
- Fe/Mg assemblage: Largely derived from volcanic/impact hydrothermal sources
- Al assemblage: Largely derived from weathering pre-deposited fine-grained glassy tephra



(Michalski et al., 2010)

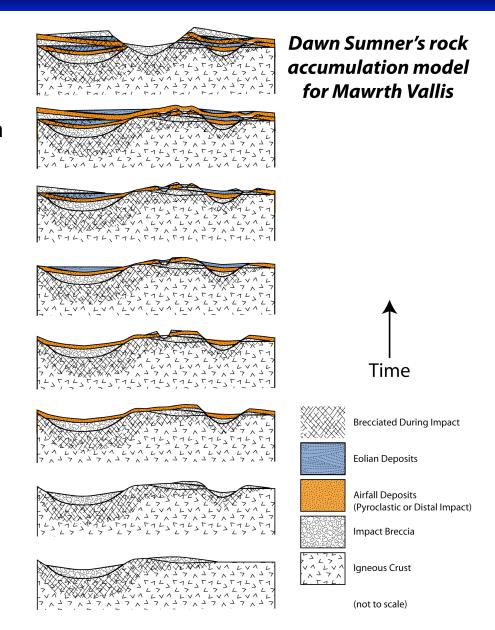
A model for rock accumulation on early Mars (esp. at Mawrth)

Heterogeneous physical processing; consistent mineralogy

- Rock accumulation primarily driven by impact processes, eolian reworking
- Predicts heterogeneous mineralogical distribution
- Continuous pyroclastic delivery & eolian reworking homogenizes mineralogy in time & space

Caveats:

- Does not exclude other modes of neoformation, but may well be dominant
- Abundances: do we understand clay production in hydrothermal & magma-volatile interaction on Mars? We may be lacking suitable analogs.



Conclusions (1)

- Weathering of pyroclastic material produces a distinct Al-rich assemblage dominated by: halloysite, imogolite, opaline SiO₂, kaolinite and lesser Alsmectite (e.g., beidellite)
- Pyroclastic delivery of clay-bearing material is a common source of Fe/Mgsmectite (and other Fe/Mg 2:1 phases) in ash fall tephra deposits on Earth
- Volcanic-hosted hydrothermal systems on Mars, or syn-eruptive alteration
 of volatile-rich pyroclastics could represent a significant flux of Fe/Mg clay
 to the early surface of Mars
- Ash transport is an effective mechanism for Fe/Mg clay contribution to the early martian sedimentary record:
 - Provide significant mineralogical overprint to complex processes
- If true, a large portion of clay mineralogy on Mars may record the spatial and temporal history of volatile-magma interaction during the earliest portion of martian history

Conclusions (2)

- This model does not require a significant role for surface water on early (Noachian) Mars. Later alteration of upper stratigraphy (presuming an Fe/Mg-rich protolith) could date to any post-depositional period.
- We view magmatic hydrothermalism (& explosive volcanism) as waning in response to decreasing magmatic volatile load and perhaps decreasing subsurface water/ice
- We view the earliest supracrustal rocks on Mars as receiving two dominant contributions to mineralogy:

1. Impacts

- Require pre-existing water to drive hydrothermalism
- May suffer from low-temperature overprinting as system "dies out"

2. Pyroclastic volcanism

- Records conditions of magmatic hydrothermalism
- Mineralogical contributions & delivery will correlate with high volatile load
- Eolian mixing will act to homogenise these two components