Earth’s earliest records of life

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Early Earth versus Mars

Questions:
• Was Mars inhabited early in its history?
• If so, are early Earth fossils and their habitats good analogues?

Answer:
• Assuming that Mars was inhabited, early Earth fossils are their best available analogue, and anything else is speculation.
This talk

• Review the types of fossil evidence for the earliest life on Earth, with some examples
• Discuss the geological settings where they occur
• Consider how they can be preserved or destroyed
• Present the sorts of prior information needed in order to find them
Problem: Old rocks are rare on Earth, mostly poorly preserved because of metamorphism, deformation and inadequate exposure, and so signs of primordial life are hard to find and even more difficult to interpret.
Early Earth

- Only a few significant Archean fossil localities
Archean Fossil Record

- Archean has a fairly diverse but quite controversial fossil record
- Doubts about age or biogenicity
Types of evidence of early life

- MICROFOSSILS: preserved remains of microbial bodies (*body fossils*)
- STROMATOLITES, ONCOLITES, MISS: preserved remains of microbial activities (*trace fossils*)
- BIOMARKERS: preserved remains of microbial molecules (*molecular fossils*)
- ISOTOPES: preserved relics of microbial metabolism (*atomic fossils*)
- BIOMINERALS: preserved remains of microbial mineralization (*no certain biominerals known from the Archean*)
- BIOALTERATION: preserved remains of microbial remineralization (*recently discovered in Archean rocks*)
Microfossils

- microbial cell walls and extracellular sheaths made up of resistant polymers
- mostly preserved where rapid entombment in microcrystalline quartz
- often not diagnostic of biological relationships as most are small and simple
- usually signs of behavior are needed before biological origin is certain
- contamination (younger fossils in older rocks) is a common problem

Proterozoic cyanobacterial microfossils
Oldest microfossils?

- ~3.45 billion years, Marble Bar, northwest Australia
- carbonaceous microstructures with apparent cellular organization
- surrounded by greenschist facies (>350°C) rocks
- pale (unmetamorphosed) in vein (contamination)?
Oldest microfossils?

- 3.23 billion years, Sulphur Springs, northwest Australia
- Pyritic filaments in subsurface hydrothermal metal deposit, oriented in different directions on substrates
- Possibly behavior but no cellular organization
Oldest microfossils?

- ~3.2 billion years, Swaziland Supergroup, South Africa
- kerogen filaments and spheres
- some possibly dividing
- but no complex behavior (e.g. mat formation, preferential orientation, colonial habit)
Oldest microfossils

• ~2.6 billion years, Transvaal Supergroup, South Africa
• dividing cells and matting filaments in stromatolites
• clear evidence of cellularity, complex behavior, interaction with surrounding environment
Stromatolites

- layered sediment mounds accreted by microbial growth, movement or metabolism
- display convex-upward doming with layers thickening over flexure crests
- microbial mats baffle currents causing grain deposition, trap sediment by overgrowth, or precipitate carbonate during photosynthesis
- modern stromatolites formed by cyanobacteria, diatoms, green algae where animals are scarce
- many abiotic mimics; e.g. stalagmites, folds

Shark Bay, Australia
Other trace fossils

- Oncolites: like stromatolites but concentric microbial lamination around a lithic nucleus

- Microbially-Induced Sedimentary Structures (MISS): surface features on sands and muds from ephemeral biofilm binding
Oldest stromatolites?

- 3.47 billion years, North Pole, Australia
- stromatolites with domical structure, layers thickening over flexure crests, alternate layers kerogen-rich and -poor, fragments eroded from apex
- very probably biogenic but absence of microfabric with palimpsest of microbes makes proving it difficult

↑1cm eroded flakes with kerogenous layering
Oldest stromatolites

- 2.72 billion years ago: clearly biogenic stromatolites in Tumbiana Formation, northwest Australia
- Show relics of tufted (middle) and erect (bottom) microbial filaments
- Implies oxygenic photosynthesis because stromatolites grew in sulfur-poor lakes on basalt, where reactants for other metabolisms are absent
Oldest trace fossils

- No certain oncolites known from Archean
- Diverse MISS known from ~3.2 Ga Moodies Group, ~3.0 Ga Pongola Group in South Africa
- Include multi-directional ripples stabilized by biofilm (*above*), eroded and fragmented superficial biofilms (*below*), gas-domes, tufted biofilm palimpsests etc
Biomarker Molecules

- hydrocarbon molecules from photosynthetic pigments and membrane rigidifiers
- found in oil and **kerogen** (non-crystalline, highly aromatic, insoluble organic matter)
- no functional groups or multiple bonds, just carbon skeleton but recognizable if complex
- concentrations minute at $T > 200^\circ C$
- easily contaminated by ancient petroleum, modern organic fluids, laboratory waste, diesel fumes

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<th>Biomolecules</th>
<th>Geomolecules</th>
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<td>Acyclic isoprenoids from pigments and Archaea</td>
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<td>Hopanepolyol and hopane from Bacteria</td>
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<tr>
<td>Sterol (cholesterol), sterane (cholestane) from Eukaryotes</td>
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Oldest molecular fossils?

- 2.69 Ga Jeerinah kerogen
- yields hopanes - Bacteria
- >C_{31} 2α-methylhopanes - photosynthetic cyanobacteria producing oxygen
- diverse steranes including cholestane – Eukaryotes
- Jochen Brocks thinks they are contaminants, but many unique features inexplicable by younger origin
- suggests all 3 domains of life were extant by late Archean, including complex eukaryotic cells
Oldest molecular fossils

- Fluid inclusions within quartz grains in the ~2.45 billion year old Matinenda Formation, Canada
- Contain oil as a fluorescing rim to the carbonic and aqueous phases filling the inclusion
- Inclusions were trapped before their host rock was metamorphosed at ~2.2 billion years ago
Oldest molecular fossils

- Matinenda fluid inclusion oil show peaks above background, so clearly indigenous
- contain diverse and abundant steranes
- $C_{27}$ steranes: contaminants in system blank; others indigenous; diversity indicates eukaryotic source
Isotopic Fractionations

- microbial metabolism fractionates the stable isotopes of carbon ($\delta^{13}C$), sulfur ($\delta^{34}S$) and nitrogen ($\delta^{15}N$)
- biotic fractionations are much larger than abiotic at low temperatures and pressures
- isotopic evidence trapped in minerals survives thermal and mechanical trauma well
- need isotopic ratios of reactant and product to be sure of biogenic origin
Carbon Isotopes

• 2 stable isotopes of carbon: $^{12}\text{C}$, $^{13}\text{C}$
• $^{12}\text{C}$ forms bonds that need slightly less energy to break
• Molecules with $^{12}\text{C}$ are slightly lighter
• These factors induce “fractionations” – discrepancies between the isotopic ratios in reactant and product
• Isotopic ratios expressed by “delta notation”

$$\delta^{13}\text{C} = \left[ \frac{\left( \frac{^{13}\text{C}}{^{12}\text{C}} \right)_{\text{sample}}}{\left( \frac{^{13}\text{C}}{^{12}\text{C}} \right)_{\text{standard}}} - 1 \right] \times 10^3 \text{[‰, PDB]}$$

• Fractionations between reactant and product expressed as “epsilon notation” e.g. $\varepsilon = \delta$ carbonate – $\delta$ organic
• “Life is lazy”, so biological reactions strongly favor the light isotope
Carbon isotope fractionations

- Different metabolisms fractionate differently.
- Fractionation range varies according to ecological conditions.
- Thus rarely diagnostic of specific metabolism.
- Can be diagnostic of autotrophic carbon-fixation if both reactant and product are preserved showing full fractionation.
C-isotope evidence of metabolism

$\delta^{13}C_{\text{organic}}$ indicates methanogenesis (maybe also methanotrophy) if $<-40\%$

- clear evidence from 2.8-2.6 billion years in non-marine Fortescue Group & marine Hamersley Group, Australia
- shows early evolution of Archaea, one of the 3 fundamental domains of life
Oldest Evidence of Life?

- ~3.85 billion years, Akilia, Greenland
- Alleged to contain graphite with light δ\(^{13}\)C (-45 to -20‰) in apatite grains in ferruginous quartzite
- Interpreted as biological isotope fractionation
- Difficult to replicate this finding
- Uncertain that rock was sedimentary
- Some doubts about age (apatite crystals hosting graphite are ~1.6 Ga)
- Without sedimentary carbonate, unclear that isotopic fractionation is of biological magnitude, or if it shows consistency expected of biology
Oldest Evidence of Life?

- ~3.8 billion years, Isua, Greenland
- graphite with light $\delta^{13}C$ (-20‰) in metamorphosed shale, consistent with biological fractionation during autotrophy
- no sedimentary carbonates (only metamorphic) to determine extent of fractionation, so can't prove biological origin

Isua meta-conglomerate showing sedimentary origin

~3.8 billion year old graphitic metasediments, Isua, Greenland
S-isotope evidence of metabolism?

- indicates sulfate reduction if $\varepsilon^{34}\text{S}_{\text{sulfate-sulfide}} > 10\%$
- clear record back to $\sim 2.7$ billion years
- single older record (3.47 billion years, North Pole)
- no evidence of hydrothermal origin (pyrite on internal growth faces of barite crystals, formerly gypsum)
- shows early evolution of respiration, but low ocean sulfate
Bioalteration

• Occurs where microbes etch or dissolve minerals
• Best known in carbonates, but also in basaltic glass, sulfides
• May form surface pits, granular texture or invasive tubules
• Tubules with complex morphology (branching, distinct opening or terminal geometry, preferential orientation towards particular minerals) can be discriminated from abiotic chemical alteration
Archean bioalteration

• Known from ~3.4 Ga pillowed basalts from Pilbara, Australia and Kaapvaal, South Africa

• Need very low metamorphic grade or else destroyed by recrystallization

• Hard to prove biogenicity unless internal organic carbon, preferential orientation towards particular minerals

Microtubules in chloritized volcanic glass, Kaapvaal Craton
Where do Archean fossils occur?

- Almost exclusively in sedimentary rocks (bioalteration in basalts are exception)
- Mostly in carbonates (silicified?), but also sulfate evaporites, clastic sediments (sandstones, shales), rarely in tuffs
- Preferentially found in fine-grained sediments
- Only known from aqueous settings (marine, lakes) but not from rivers, ephemeral water bodies
- Largely preserved in reducing environments
- Stable tectonic settings distal from point sources of heat or pressure (volcanoes, orogenies or impact craters)
- In areas of active erosion
What do you need to find them?

• Thorough mineralogical and lithological information
• Full stratigraphic history
• Good geochronology
• Complete post-depositional paragenesis (diagenesis, metamorphism, deformation)
• Recent geomorphological evolution
• A diverse suite of analytical techniques (macro-imaging, microscopy, whole-rock geochemistry, organic geochemistry, isotopic geochemistry)
• A whole lot of luck!!
Where would I go to find them on Mars?

- Somewhere with diverse sedimentary rocks
- Somewhere with long-lived water bodies
- Somewhere with rapid lithification (evaporative or hydrothermal settings)
- Somewhere with fine grain-size
- Somewhere with little evidence of oxidation or acidic alteration
- Somewhere with little subsequent disturbance
- Somewhere only recently exposed
- Somewhere with well-understood basic geology
Early Mars?
• Nakhla
Nitrogen isotopes

- Positively fractionated $\delta^{15}N$ in organic matter by denitrification
- No strong evidence until early Proterozoic, perhaps in late Archean
Ion microprobe (CAMECA imf1270) carbon isotope analyses
Ueno et al. 2001, IGR 43, 196-212

Nitrogen isotopic composition of organic matter through geological time

- Modern $n = 4$
- Precambrian $n = 1$

- Modern <0.5 Ga.
  Peters et al., 1978
- Late Proterozoic
  0.7 to 0.9 Ga.
- Early Proterozoic
  1.6 to 2.1 Ga.
- Late Archean
  2.5 to 2.7 Ga.
- Early Archean
  3.4 to 3.5 Ga.

- $\delta^{15}N_{\text{AIR}}$ %
  -50
- T.O.C.
  -37.2 %

Reductive acetyl-CoA pathway
Calvin cycle
Thermoproteus neutrophilus
Oscillatoria williamisi
Synechococcus sp.
3-hydroxypropionate cycle

Desulfbacterium autotrophicum
Methanobacterium thermooautotrophicum
Thiobacillus neapolitanus
Thiomicrospira sp.
Chloroflexus aurantisius

Thermoproteus neutrophilus
Oscillatoria williamisi
Synechococcus sp.
3-hydroxypropionate cycle
• Pyrite corrosion
  Sulfolobus
Oldest evidence of life?

- Coonterunah Group, northwest Australia - 3.52 billion years
- Oldest low-grade (sub-amphibolite facies) rocks on Earth
- Mostly minimally deformed (only tilting and block faulting)
- Unconformably underlies the Warrawoona Group, previously thought to be the oldest low-grade rocks on Earth
- Therefore a good place to go looking for C-isotope evidence of autotrophy
Coonterunah Group

- 5km of stratigraphy
- 75km strike length
- Units mostly continuous
- Broad regional folding
- Metamorphic grade increases into western fold and towards granite
- Deep-water marine depositional environment (no current structures)
- Distal to land (no coarse terrigenous clastic sediments)
Coonterunah basalts

- Bulk of Coonterunah stratigraphy is basalt
- Mostly tholeiitic, but magnesian and komatiitic in places
- Mostly massive, but extensive pillowed horizons and some amygdaloidal units
- Some felsic volcanics (rhyolites & dacites) interbedded, massive or autobrecciated

Pillowed magnesian basalt from base of Coonterunah Group
Coonterunah tuffs

- Several metres of tuff underlie carbonate
- Conglomeratic at base, fining upwards to mud
- Basaltic in composition
- Some carbonate matrix to conglomerate, but elsewhere slightly silicified
- Indicates that carbonate was not formed after deposition
Coonterunah carbonates

- One carbonate horizon
- Only 32 cm thick
- Can be traced along strike for 5 km
- Plane-laminated at 1 mm scale
- Inter-laminae of magnetite
- Composed of ferroan calcite
- No detrital component
- No sedimentary structures indicating current activity
- Not metamorphic or metasomatic

Pure ferroan calcite

Ferroan calcite interlayered with magnetite
Coonterunah chert

- Chert overlies carbonate bed and comprises bulk of Coonterunah sediment
- Alternating 1 cm laminae of microcrystalline quartz and magnetite-grunerite-kerogen
- Rare siderite and pyrite laminae
- No traction current structures
- Resembles Fe-poor banded iron formation

Magnetite – quartz laminated chert
Oldest Evidence of Life?

- Cant say what type of C-fixation (oxygenic photosynthesis, anoxygenic photosynthesis, chemoautotrophy etc.) because fractionation is not diagnostically light or heavy
- But definitely evidence of life!
- Could be argued that this is oldest indisputable evidence for life on Earth (and in the Universe)

- Older light $\delta^{13}C_{\text{organic}}$ from $\sim3.8$ Ga rocks at Akilia and Isua, Greenland don’t have co-existing sedimentary $\delta^{13}C_{\text{carb}}$
- Coonterunah data more robust because isotopes on both sides of the carbon-fixation reaction are preserved
Oldest molecular fossils

- Matinenda biomarkers show peaks above background, so clearly indigenous
- Abundant $>C_{31}$ 2α-methylhopanes are derived from cyanobacteria

![Graph showing Matinenda oil inclusions with peaks labeled $C_{31}2\alpha$(Me), $C_{30}\alpha\beta$, $C_{32}2\alpha$(Me), and $C_{31}\alpha\beta$. The graph also shows Laboratory system blank and Final outside rinse.]
Life in the Archean

- Can place dates on the “Tree of Life” based on the Archean fossil record
- Shows most microbial evolution occurred early including Eukaryote lineage
- Earth has been inhabited since almost the start of the geologic record
- Life probably appeared rapidly – didn’t take billions of years
- May mean that it evolves easily given the right conditions
- Widely distributed in the Universe?